

Supplementary information

Hallis et al. (2025) The D/H ratio of the Martian Mantle: Evidence from the Lafayette Meteorite. *Advances in Geochemistry and Cosmochemistry* 1(2): 761, <https://doi.org/10.33063/agc.v1i2.761>

S1. Hydrogen isotope measurements – additional information.

The amount of H₂O background contamination within the analytical chamber was determined via analysis of the nominally anhydrous San Carlos olivine standard, which was analysed twice each day throughout the analytical session. The average background value was subtracted from all standards and unknown phases analysed in this study; the calibration line was corrected for this H₂O background and forced through the origin. The H₂O background within the analytical chamber averaged ~7 µg/g. However, since the Lafayette thin-section was not co-mounted with the standards, it was not possible to estimate the background on the sample in this manner. Instead, a background of approximately 263 µg/g was estimated via analysis of pyroxene in this sample assuming that the H⁺/O⁻ calibration remained consistent between mounts. The higher H background most likely results from both the epoxy mounted nature of the sample producing H contamination and a real water signature present within the Lafayette pyroxene. Hence, this value serves as a maximum background value, which is in any case more than an order of magnitude lower than the Lafayette apatite with lowest H₂O abundance. Based on repeated analyses of the Durango apatite standard throughout each analytical session, errors from counting

statistics, and uncertainty in the instrumental background, we estimate the external reproducibility (2σ SD) of the H_2O concentrations presented here to be on average $\pm 15\%$, similar to Mane et al. (2016), Stephan et al. (2018) and Davidson et al. (2020).

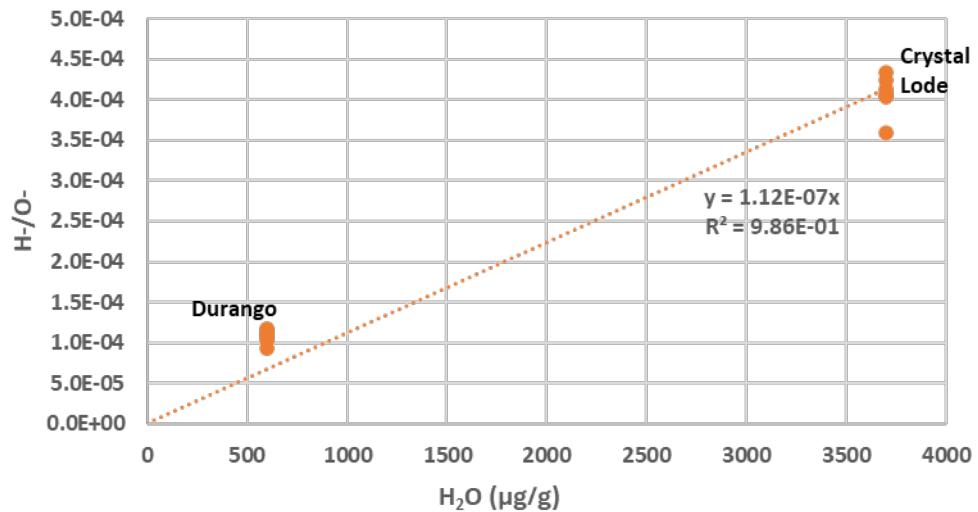


Figure S1: $\text{H}_2\text{O}/\text{H}_2\text{O}^-$ vs. H_2O ($\mu\text{g/g}$) calibration line calculated using Durango and Crystal lode standard apatite crystals. This calibration line is background corrected and forced through the origin.

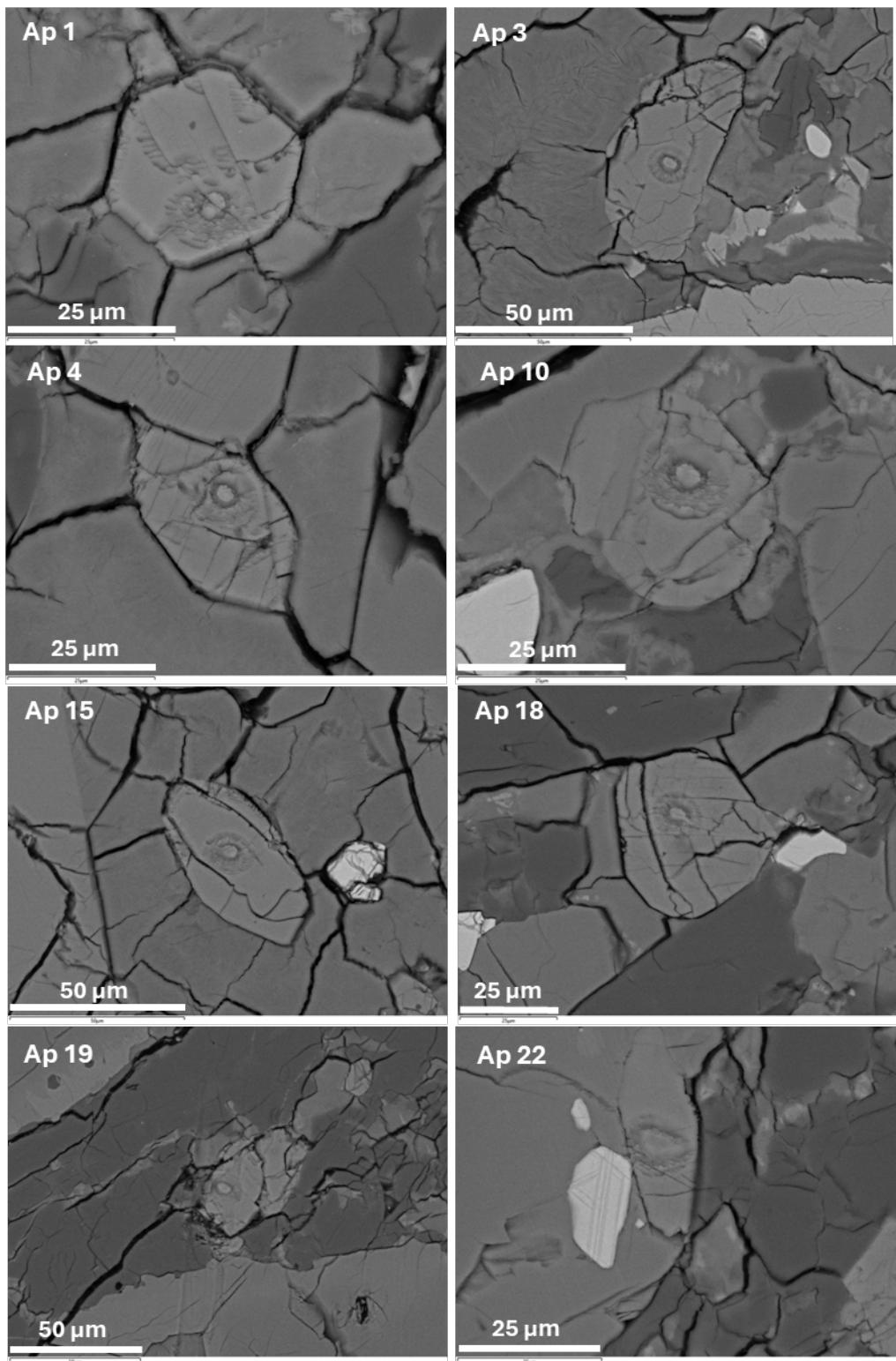


Figure S2: Backscatter electron images of the Lafayette apatite SIMS pits.

Table S1: Electron microprobe uncertainties and detection limits for each measured element (wt %)

| Element | Uncertainty (2 σ)* | Detection limit ($\mu\text{g/g}$) |
|---------|----------------------------|-------------------------------------|
| P | 0.328 | 913 |
| Si | 0.017 | 57 |
| Ti | 0.005 | 131 |
| Al | 0.004 | 57 |
| Cr | 0.024 | 81 |
| Fe | 0.020 | 124 |
| Mn | 0.009 | 122 |
| Mg | 0.002 | 79 |
| Ca | 0.270 | 669 |
| Na | 0.023 | 805 |
| K | 0.007 | 62 |
| S | 0.022 | 56 |
| F | 0.216 | 1972 |
| Cl | 0.032 | 57 |

*n=3 analyses of standard apatite

Supplementary References

Davidson J., Wadhwa M., Hervig R.L. and Stephan A. (2020). Water on Mars: Insights

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Mane P., Hervig R., Wadhwa M., Garvie L. A. J., Balta J. B. and McSween Jr H. Y. (2016).

Hydrogen isotopic composition of the Martian mantle inferred from the newest

Martian meteorite fall, Tissint. *Meteoritics and Planetary Science* 51, 2073-2091.

<https://doi.org/10.1111/maps.12717>

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<https://doi.org/10.1038/s41598-018-30807-w>